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DETERMINATION OF LITHOLOGIC BELONGING AND RESERVOIR PROPERTIES OF CLASTIC FORMATIONS USING FIELD GEOPHYSICAL SURVEY DATA OF BAKLANOVSKOE FIELD

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УСТАНОВЛЕНИЕ ЛИТОЛОГИЧЕСКОЙ ПРИНАДЛЕЖНОСТИ И КОЛЛЕКТОРСКИХ СВОЙСТВ ТЕРРИГЕННЫХ ПЛАСТОВ ПО ПРОМЫСЛОВО-ГЕОФИЗИЧЕСКИМ ДАННЫМ БАКЛАНОВСКОГО МЕСТОРОЖДЕНИЯ

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The paper contains results of comprehensive studies of well logging data and laboratory study of the core. Studies were carried out to determine lithologic composition and reservoir properties of clastic formations of Baklanovskoe field. It is noted that grains that compose clastic rocks are diversified in size, shape, degree of roundness and sorting, granulometric and mineral composition, type and composition of cement, its structure and ratio with grains. It is stated that these features of rocks determine structure of pore space, as an ordered set of mutual relations of elements of different hierarchical levels (for example, mineral grains, samples and rock formations). It is noted that study of pore space structure, reservoir properties and lithologic features is performed based on fractional composition of rocks. However, it is impossible to understand in details entire section of even exploration and prospecting wells in case of poor core recovery. Therefore layer-by-layer description (including wells with no core) is carried out based on logging linked to results of core analysis. It was found that logs can be used for lithologic distinguishing and selection of lithologic series of clastic rocks with its material composition determination. Information about lithologic series also allows approaching to a solution for the problem of cyclic sedimentation. Equations of multiple correlations between reservoir properties of rocks and amount of psammite, silty and pelitic particles in rock matrix. It is possible to determine porosity by recorded natural radioactivity on gamma ray charts. Solving the above mentioned problems is considered on the basis of a systematic structural approach.

Ключевые слова:

геофизические исследования скважин, керн, литология, коллекторские свойства, фракционный состав, структура порового пространства, уравнение регрессии, корреляционные связи, системный подход, гамма-каротаж, физические свойства, геологическая неоднородность, гистограммы, радиоактивность, литоряд, пласт-коллектор.

Изложены результаты комплексных исследований материалов геофизических исследований скважин и лабораторного изучения керна при определении литологического состава и коллекторских свойств терригенных пластов Баклановского месторождения. Отмечено, что зерна кластического материала, слагающего терригенные породы, разнообразны по своим размерам, форме, степени окатанности и отсортированности, гранулометрическому и минеральному составу, по типу и составу цемента, его строению и соотношению с зернами. Указывается, что эти особенности пород определяют структуру порового пространства как упорядоченную картину взаимных отношений элементов различных иерархических уровней (например, минеральных зерен, образцов и пластов горных пород). Отмечено, что изучение структуры порового пространства, емкостно-фильтрационных свойств и литологических особенностей осуществляется по данным о фракционном составе пород. Однако при недостаточном выносе kernового материала даже из разведочных и поисковых скважин невозможно детально представить весь разрез скважины. Поэтому послынное описание разрезов скважин (и в том числе бескерновых) осуществляется на основе геофизических исследований скважин, связанных с kernовыми данными. Установлено, что по каротажным диаграммам можно проводить не только литологическое расчленение разреза, но и выделение литорядов терригенных пород с определением их вещественного состава. Сведения о литорядах позволяют также подойти к решению задачи о цикличности осадконакопления. Получены уравнения множественной корреляции между коллекторскими свойствами продуктивных пород и содержанием псаммитовых, алевритовых и пелитовых частиц в скелете породы. Показана возможность определения пористости по зафиксированным показаниям естественной радиоактивности на диаграммах гамма-каротажа. Решения вышеперечисленных задач рассмотрены на базе системно-структурного подхода.

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Introduction

Study of relations between parameters that characterize any geological formation is a question of interest in oil industry. Depending on conditions of their formation and distribution, rocks have essential structural and textural features.

In section of clastic deposits rock formations are characterized by a certain set of physical properties – porosity, permeability, clayiness, water saturation, density, elasticity, specific electrical resistance, radioactivity etc.

Lithological (granulometric composition, type of cement, presence of accessory minerals) and geometric (thickness, depth, occupied area) properties of rocks also play a significant role.

These properties are studied on the basis of well geophysical and field geological studies and laboratory studies of rock samples. Each of mentioned studies characterizes individual elements of geological formations that often belong to different hierarchical levels of the geological system.

For the first time a question of systematic study of natural formations was raised in early 30's by famous biologist L. von Bertalanffy [1], who formulated the basic theses of a theory of system studies, which further development is connected to works of A.A. Liapunov, A.A. Malinovskii, D. Nidkhem, A. Rapoport, V.N. Sadovskii, Iu.A. Urmantsev, U.R. Eshbi and others. Questions of application of system approach in geology are considered in works of Iu.A. Voronin, L.F. Dementev, A.B. Kazhdan, Iu.N. Karagodin, L.D. Knoring, Iu.A. Kosygin, A.I. Kholin, I.P. Sharapov, Iu.V. Shurubor and several other researchers [2-7].

A system approach considers researchers should show that geological formation being under study can be considered as an interconnected multilevel system, and find out elements it consists of. So, for example, clastic deposits of oil and gas fields that belong to Visean age of Perm Prikamye are presented mainly by sandstones, siltstones and mudstones. The grains of clastic material composing these rocks are diverse in its size, shape, degree of roundness and sorting, granulometric and mineral composition. In addition, sand-clayish rocks also differ in composition and type of cement, its structure

and ratio with grains. All these features of rocks determine structure of pore space, which represent position of elements in space order (for example, grains, mineral particles in rock matrix), of an integral geological formation (sample, layer) associated with nature of these properties. It should be noted that if granulometric composition is a characteristic of rock structure, then for a structure of pore space of clastic reservoirs such a characteristic is volume of interconnected pores and pore channels.

When determining its structure a geological formation itself should be considered as a natural system consisting of elementary parts. The total of system elements allows determination geometry of void space of rocks. Structural relations are important only in context where they characterize stability of a system. Therefore, a different combination of structural characteristics of rocks determines a degree of stability of a system and its heterogeneity [8–11].

So, due to cementation of particles, pore channels narrow down and even their closure, which often leads to formation of dead-end zones. All this complicates a structure of pore space of rocks. Distribution of grains in rock volume and their granulometric coefficients influence formation of void space strongly [3, 12–15].

The relationship between structure of rock matrix and structure of pore space is most clearly presented on example of stacking and cementation of particles. Heterogeneity of production formations (as a degree of stability of geological formation system) cause a difference in their reservoir properties.

Reservoir properties (porosity, permeability, oil saturation, clayiness, residual water content) of individual reservoirs and fields, belonging even to the same tectonic structure, of the same origin and sedimentation conditions have significant differences.

It should be noted that there are certain regularities between mentioned reservoir parameters that represent properties of geological formation at higher level of hierarchy (sample, formation), which is logically related to characteristics of a formation at lower (structural) level (grains, pores) [3, 5].

To describe structural features of geological formations of lower level it is needed to use

coefficients of granulometric studies that are as follows: average grain size (median diameter Md), coefficient of asymmetry K_s and coefficient of heterogeneity S . These coefficients reflect the main features and nature of distribution of grains as a building material involved in formation of rock and void space [3, 16].

In turn, variability of production reservoir parameters and degree of their geological heterogeneity (i.e. macrostructure of pore space) as properties of geological objects at a higher structural level can be characterized by coefficients of sandiness k_{sand} , stratification factor k_{strat} , connectedness k_{con} , effect k_e etc. [3].

Fractional composition is one of the main characteristics of clastic rocks. A structure of pore space and thus reservoir properties and lithological features of rocks depend on it. Fractional composition is determined by core samples in laboratory conditions. However, core material is usually not enough to represent in detail entire well section. It characterizes individual intervals of well section only [10, 17]. Therefore, conventional formation of layer-by-layer description of well section is carried out on the basis of continuous displaying of well section by well logging diagrams associated with core data. As is known, when describing a core, a lot of lithological differences of rocks are recorded, which cannot be identified by well logging curves. Only rocks with sharply outlined lithological and physical properties such as clay rocks, siltstones, sandstones etc., i.e. such groups (lithological series), which combine similar in material composition and physical properties of rocks variety can be distinguished by logs. Usually, laboratory description of core of clastic deposits is based on results of granulometric and instrumental analysis of thin sections.

Determination of lithology and reservoir properties of rocks based on results of core treatment

Most often lithological and structural properties of clastic rocks are determined from granulometric composition data, nature of its packaging, content and composition of cement in rock matrix. Reservoir properties are determined by gas saturation and flow method.

Thus, open porosity of clastic rocks of Visean age of Baklanovskoe field varies from 2.7 to 27.8 %. A section of this field is represented mainly by sandstones, silty sandstones, silt and clay rocks. Carbonate rocks (limestones of varying degree of dolomitization) play a subordinate role in a roof of Tula layer.

Sandstones are quartz, light gray and gray, poorly cemented, with carbonaceous inclusions, carbonized plant remains and inclusions of pyrite. Sandstones are heterogeneous, with various sorting and rounding of mineral grains. Content of pelite fraction C_{PI} (smaller than 0.01 mm) vary from 3.3 to 14.4 %. A prevalence of psammitic fraction of C_{PS} with a particle size of 0.25-0.1 mm in sandstones is 47-82 %. The content of coarse-grained fraction (more than 0.5 mm) does not exceed 0.3 %. Fractions with size of 0.5-0.25 mm rare make up 10 %. Content of siltstone fraction of C_{AI} (0.1-0.01 mm) varies from 9.4 to 23.2 %. Cement in sandstones is more often clayish and carbonaceous-clayish, less often calcite, carbonate and mixed. Sometimes cementation of rock matrix is done due to consolidation of grains.

Siltstones have different grain size. They are quartz, gray, light and dark gray, carbonaceous and carbonaceous-argillaceous, layered. Siltstones are composed of detrital material, usually have worse roundness than sandstones. Content of carbonate cement C_C in rock matrix (as in sandstones and silty sandstones) is usually small. In general, the silt fraction predominates (up to 66 %), content of fractions with particle size of 0.25-0.1 mm is 22-23 %.

Silty sandstones can be found in section rarely. According to granulometry data they occupy an intermediate position between sandstones and siltstones (Table 1, Fig. 1). It is indicative that silty sandstones have the same content of psammitic and silty fractions (42-47 % each).

The argillites are predominantly dark gray, almost black, unevenly silt and carbonaceous, thin-layered and slate. Content of C_{PI} is more than 50 %, content of C_{AI} is about 40 %.

Sandstones, silty sandstones and siltstones refer to oil-bearing rocks (see Table 1). In terms of gas permeability they are divided into four classes: I – unproductive reservoirs ($K_{perm} < 1$ mD), II – low-productive ($K_{perm} = 1-10$ mD), III – medium-productive ($K_{perm} = 10-160$ mD), IV – highly productive ($K_{perm} > 160$ mD).

Table 1

Characterization of oil-bearing rocks by permeability and granulometric composition

Oil-bearing rock	K_{perm} , mD	Granulometric composition, %					
		psammites			silt	pelites	carbonate particle
		> 0,5 mm	0,5–0,25 mm	0,25–0,1 mm	0,1–0,01 mm	< 0,01 mm	
Sandstones	< 1	0,16	13,81	47,03	18,28	14,40	6,32
	1–10	0,21	4,04	57,33	23,11	12,86	2,45
	10–160	0,07	3,27	64,93	23,18	7,78	0,77
	> 160	0,24	4,73	82,23	9,37	3,26	0,17
Silty sandstones	< 1	0,07	0,43	42,33	43,00	14,17	–
	1–10	0,03	0,50	43,73	44,17	11,60	–
	10–160	0,08	1,52	44,84	46,50	6,28	0,78
	> 160	0,10	1,86	45,90	47,00	5,00	0,14
Siltstones	< 1	0,14	1,09	22,47	59,70	13,69	2,91
	1–10	0,09	0,95	23,36	61,40	12,81	1,39
	10–160	0,12	1,07	23,23	65,99	8,00	0,96
	> 160	–	1,02	31,35	56,50	10,48	0,65

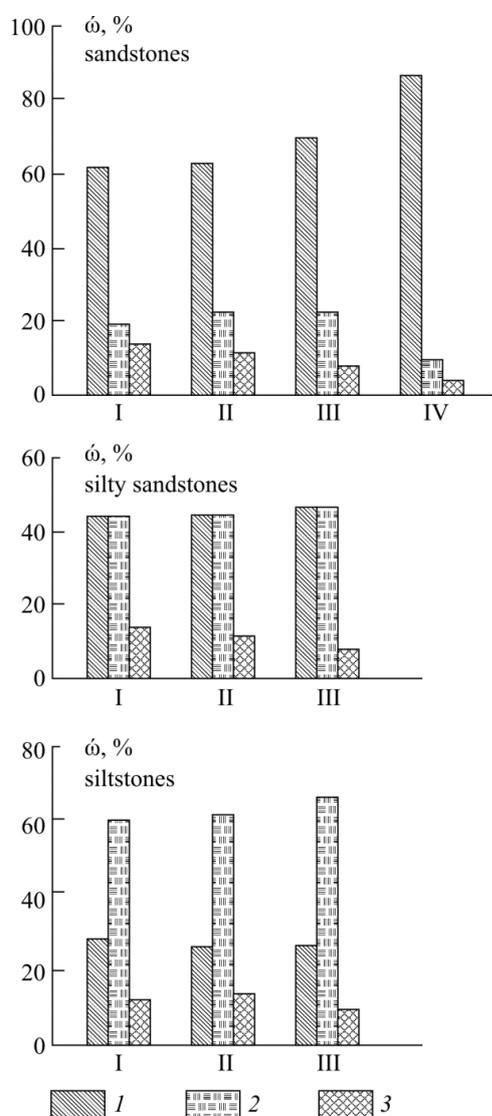


Fig. 1. Histograms of fractional composition of Tula clastic deposits of Baklanovskoe field: ω – fraction content, %; 1 – sandstones, 2 – silty sandstones, 3 – siltstones; I < 1 mD; II – 1–0 mD; III – 10–160 mD; IV > 160 mD

Analysis of the data of Table 1 and Fig. 1, 2 shows that content of pelitic and psammitic fractions has the greatest and decisive influence on flow properties of rocks (permeability). The same effect have these fractions on volume capacity (porosity).

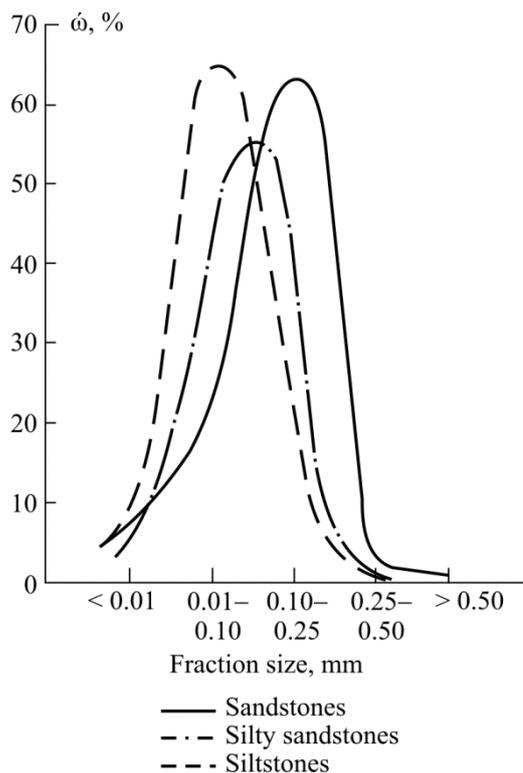


Fig. 2. Fractional composition distribution curves in different groups of rocks in class of medium-productive reservoirs ($K_{perm} = 10-160$ mD)

Content of carbonate material for different classes of reservoirs in terms of permeability vary spontaneously and unsystematically and does not always correspond to certain lithological variety of clastic rocks.

This fact is confirmed by equations of multiple correlation between coefficients of porosity and permeability and content of psammitic, silty, pelitic fractions and content of carbonate particles for clastic reservoirs of Baklanovskoe field:

$$K_p = 0.45 C_{Ps} + 0.4 C_{Al} + 0.19 C_{Pl} - 24.6, \quad (1)$$

$$K_{perm} = 8.93 C_{Ps} - 20.71 C_{Pl} + 70.35, \quad (2)$$

where K_p and C , %, K_{perm} , mD.

The equation (1) shows the capacitive properties of reservoirs are estimated by content of all fractions. Flow properties are determined mainly by values of pelite fraction (2), i.e. in fact, according to content of clay particles in matrix of rock. Insignificant content of carbonate particles does not have a noticeable effect on reservoir properties, which is confirmed by absence of correlation with C_C values in both equations.

It should be noted that in addition to using results of granulometric analysis of core to determine porosity and permeability in laboratory conditions, it is compulsory to study irreducible water saturation of rocks. It is known that in reservoir conditions productive rocks contain a certain amount of irreducible water. Irreducible water saturation of clastic rocks is usually determined by indirect laboratory methods such as capillary pressure and centrifugation.

It was found, for example, that irreducible water saturation of $K_{i.w}$ of clastic reservoirs of Lower Carboniferous of Baklanovskoe field varies widely, in particular from 16 to 57 % for sandstones and from 28 to 96 % for siltstones. According to these data, siltstones contain irreducible water in much larger quantities than sandstones.

While considering $K_{i.w} = f(K_p)$ и $K_{i.w} = f(\lg K_{perm})$ there is a strong trend of irreducible water amount to decrease with increase of porosity and permeability (Fig. 3, a, b).

Regression equation $K_{i.w} = f(K_p)$ is characterized by a high density of relationships between compared parameters correlation coefficient $r = 0.76$) and has following analytical equation:

$$K_{i.w} = 114.8 - 0.56 K_p^2 + 1.5 \cdot 10^{-2} K_p^3. \quad (3)$$

There is even higher dependency between irreducible water saturation and permeability ($r = 0.87$):

$$K_{i.w} = 95.03 - 52.3 \lg K_{perm} + 7.67 (\lg K_{perm})^2. \quad (4)$$

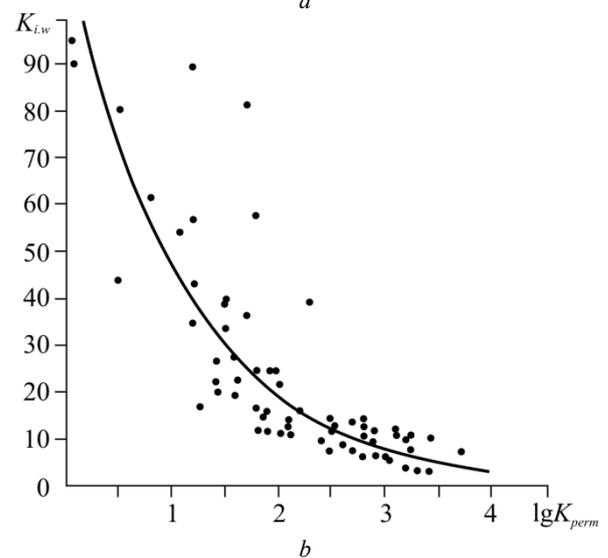
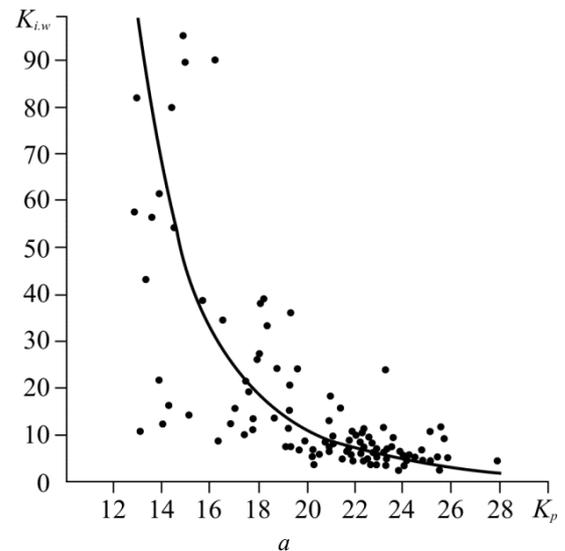


Fig. 3. A function of irreducible water saturation of clastic reservoirs from: a – K_p ; b – $\lg K_{perm}$

According to results of laboratory studies, it is established that reservoirs with porosity of less than 14 % contain more than 70 % of irreducible water.

$$K_{i.w} = 114.8 - 0.56 K_p^2 + 1.5 \cdot 10^{-2} K_p^3. \quad (5)$$

An even closer relationship is observed between irreducible water saturation and permeability ($r = 0.87$):

$$K_{i.w} = 95.03 - 52.3 \lg K_{perm} + 7.67 (\lg K_{perm})^2. \quad (6)$$

According to results of laboratory studies, it was established that productive rocks with porosity of less than 14 % contain more than 70 % of irreducible water.

The highest value of irreducible water saturation is caused by high content of clay

particles in cement of rock, which, in turn, are capable of retaining water in significant quantities due to their large adsorption surface. Between irreducible water saturation $K_{i,w}$ and clayiness K_{clay} , there is a rectilinear dependence with a high correlation coefficient ($r = 0.89$):

$$K_{i,w} = 4.05 K_{clay} - 2.85. \quad (7)$$

Determination of lithological composition and reservoir properties of rocks according to well logging

For complete characterization of clastic rocks by content of fractions gamma-ray (GR) logging method is promising. GR-method registers natural radioactivity of rocks, caused mainly by content of thorium (Th), radium (Ra) and radioactive potassium isotope (K^{40}) in it [18].

Study of spectral composition of natural gamma radiation of reservoirs represented by quartz sandstones in oil fields in Perm region made it possible to reveal patterns of distribution of radioactive elements (Th, Ra, K^{40}) depending on granulometric composition of rocks. It was found that total content of radioactive elements relative to pelitic, psammitic and silty fractions is characterized by high tightness of bond, but their character for each fraction is different. This indicates a possibility of establishing dependencies between different fractions and GR readings.

Based on numerous studies, it was established that application of well logging methods, such as potential of spontaneous polarization (PS) and GR, is legitimate during determination of reservoir properties in clastic reservoirs [8, 11, 19-21].

In our particular case, the most informative method for determining permeability in clastic reservoirs is GR method. Application of PS method is limited due to weak differentiation of PS curve in wells drilled in salt drilling fluids. Wells with good quality of recorded PS are very few.

As is known, GR is used to study geological sections of wells and is based on differentiation (separation) of rocks from their natural γ -activity. The essence of GR is to study natural gamma field along a wellbore by recording intensity of radiation that occurs during spontaneous decay of radioactive elements in rocks [10, 19, 22]. Geiger-Muller or more efficient scintillation counters are used as an indicator for section determination. Obtained during measurement curve characterizes

intensity of gamma radiation of formations along a wellbore is called gamma-ray logging curve.

Natural radioactivity of rocks is caused mainly by presence of natural radioactive elements – uranium U and its decay products, radium Ra, thorium Th and radioactive isotope of potassium ^{40}K . Remaining radioactive elements such as rubidium, samarium, lanthanum, lutetium etc. do not make a significant contribution to total natural radioactivity of rocks.

Radioactivity of sedimentary rocks (a subject of our research) is directly dependent on degree of their clay content. Clay sandstones and siltstones are characterized by intermediate values of radioactivity i.e. between radioactivity of pure rocks (e.g. clean sandstones) and radioactivity of clay rocks. Enrichment of sedimentary rocks with radioactive elements occurs due to their precipitation or adsorption by finely dispersed and colloidal particles. Therefore, due to large specific surface, clay rocks (clays, argillites, clay siltstones) sorb a large number of radioactive elements during sedimentation process and are characterized by increased indications in GR diagrams. Besides, a possibility of extensive use of GR method for determining reservoir properties of rocks is that the GR method represents certain extent geometry of void space and reveals stable relationship with fractional composition of rocks, since the structure of pore space of clastic rocks is directly related to reservoir properties. It should be noted that radioactivity of carbonate rocks is usually 1.5-2.0 times lower than that of sandy clay rocks [19, 22].

Study of cliff section is of particular interest, since clay parameter determines reservoir properties of rocks, in particular their porosity and permeability [8, 9, 23].

So, natural radioactivity of rocks basically depends on mineral composition and cement of rocks. Registered radiation J_{γ}^{res} in front of formation is sum of emissions of a formation, washing liquid and background of device. In general, intensity of natural radioactivity detected by wellbore radiometer (GR method) is directly proportional to radioactivity of rocks traversed by a well [9, 10, 19].

Since magnitude of radioactivity of sedimentary rocks correlates well with their clay content, it is possible to pick formations with different content of clay in rock matrix in clastic section on GR diagrams. Configurations

of GR curves during logging are usually distorted due to presence of an integrating cell, which causes inertia of apparatus [24]. Therefore, GR diagrams appear asymmetric with respect to middle of formation and shifted in direction of movement of device. Boundaries of reservoir (with reduced radioactivity) are determined from points corresponding to onset of decay of GR curve in base of formation and beginning to rise at its roof [25].

On the one hand the basis of a method for determining porosity by GR is correlation between porosity of clastic rocks and clayiness of $K_p = f(C_{clay})$, on the other hand between clayiness and natural radioactivity of the rocks, $\Delta J_\gamma = f(C_{clay})$.

GR-curves registered in wells with different measurement conditions and radiometric equipment features are not comparable in their quantitative interpretation. In order to exclude influence of a source of neutrons in NGR channel, design features of measuring equipment, background and well conditions on GR values relative values of gamma activity of reservoir (double difference parameter ΔJ_γ) (Fig. 4) is used. Usually, reference formations are dense limestones of Tornaisian stage with minimum GR values ($J_{\gamma \min}$) and clays of Tula layer with maximum GR values ($J_{\gamma \max}$) [10].

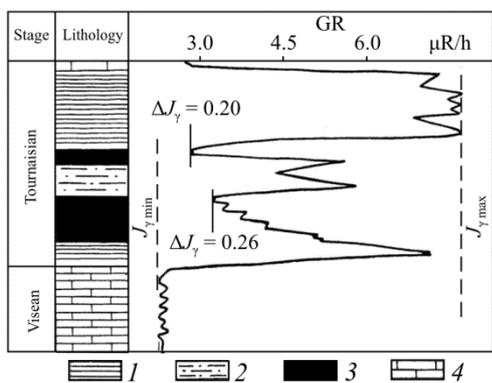


Fig. 4. Calculation of ΔJ_γ by GR: 1 – clay; 2 – siltstones; 3 – reservoir; 4 – limestone

Parameter ΔJ_γ is calculated by equation

$$\Delta J_\gamma = \frac{(J_\gamma^{\text{res}} - J_{\gamma \min}) \pm \delta J_\gamma}{J_{\gamma \max} - J_{\gamma \min}}, \quad (8)$$

where J_γ^{res} – value of GR in front of a reservoir; $J_{\gamma \max}$ – maximum value of GR in front of clays; $J_{\gamma \min}$ – minimum value of GR in front of tight limestones; δJ_γ – correction that consider changes in recorded intensity of gamma ray as a

function of speed of movement of tool V , constant of time of integrating cell Δt and formation thickness h . Corrections are introduced for low thickness formations according to an equation $h \geq 4Vt / 3600$.

In order to determine K_p from GR for studied field a function $\Delta J_\gamma = f(K_p)$ is used. For instance, established dependence for clastic reservoirs of Baklanovskoe field is characterized by a high correlation coefficient ($r = 0.92$) and has following analytical equation:

$$K_p = -69.7\Delta J_\gamma^3 + 96.3\Delta J_\gamma^2 - 63.9\Delta J_\gamma + 25.2. \quad (9)$$

Together with other methods of well logging GR data is also used for definition of lithologic well section and its correlation, for definition of reservoirs and its clayiness evaluation. Under favorable conditions GR materials are used for indirect determination of porosity, irreducible water saturation and permeability of reservoir rocks [8, 10].

GR records (parameter ΔJ_γ) can also be used for express determination of clastic rock lithologic sequences for a specific field. As an example, double difference parameter ΔJ_γ as a function of fraction content were built and analyzed for Baklanovskoe field (Fig. 5).

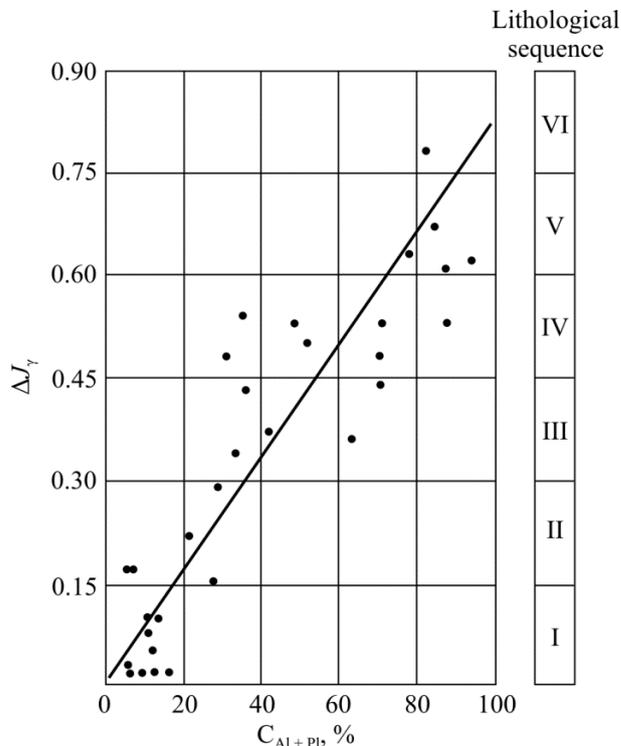


Fig. 5. Function $\Delta J_\gamma = f(C_{Pl + Al})$ and lithological sequences of rocks

The most tight relation (correlation coefficient $r = 0.91$) belongs to function of total saturation of pelite and silt $\Delta J_\gamma = f(C_{PI + AI})$. The equation of regression is as follows:

$$\Delta J_\gamma = 8.2 \cdot 10^{-3} C_{PI + AI} + 0.01. \quad (10)$$

It was found that content of clay particles for each core sample is about one-third of total content of silt and clay particles. This fact allows giving a fairly real description of lithologic sequences of clastic rocks that occur in sections of wells. All clastic rocks were divided

into six lithologic sequences: three in oil-bearing intervals and the same in enclosing rocks (non-reservoirs). An interval with ΔJ_γ equal to 0.15 was chosen as a basis. Such choice is conditioned by the fact that among oil-bearing formation traditionally there are three groups of reservoirs with low, medium and high reservoir properties. Herewith lower limit of porosity ΔJ_γ is 0.45. According to chosen interval of ΔJ_γ values lithologic sequences were picked up even in non-oil-bearing part of well section (Table 2).

Table 2

Characteristics of lithologic sequences, picked up from GR diagrams

Lithologic sequences	Interval average porosity, %	Interval average fraction content, %			Content of lithologic sequence	ΔJ_γ , fractions
		Ps	AI	PI		
I	22.1	92.1	5.3	2.6	Clean sandstone, silty sandstone	< 0.15
II	16.5	73.7	17.6	8.7	Silty sandstone, sand silt, sandy siltstone	0.15–0.30
III	12.5	55.3	30.0	14.7	Siltstone-clayish sandstone, siltstone	0.30–0.45
IV	10.4	37.0	42.2	20.8	Clayish sandstone, clayish-sandy siltstone	0.45–0.60
V	–	17.3	55.5	27.5	Clayish siltstone, sandy mudstone	0.60–0.75
VI	–	< 8.0	62.0	30.0	Very clayish siltstone, clay rocks, mudstone	> 0.75

Conclusion

Results of conducted studies on core of Baklanovskoe field allow to differentiate structure of pore space of reservoirs on basis of fractional composition data of rock samples. A close relationship between reservoir properties of oil-bearing rocks and content of psammite, silty and pelitic fractions is observed.

Oil-bearing rocks are characterized in terms of permeability and granulometric composition. Reliability of received information is confirmed by

equations of multiple correlation with high correlation coefficients. It is shown that there is a clear tendency of irreducible water saturation to decrease with increase in reservoir properties of formation.

It is found that GR logging is the most promising method among spectral methods of core analysis allowing both to determine porosity of oil-bearing formations in wells with no core and pick up lithological sequences in oil-bearing and non-oil-bearing areas of well section. It helps to reveal complete lithological and structure characteristics of rocks.

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